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SEISMIC HAZARD SMULATION AND PREDICTION

Paul Somerville URS Corporation Pasadena

OUTLINE

Ground Motion Needs for Performance Based Engineering

Seismic Hazard Analysis Procedure

Ground Motion Simulation and Prediction

•Source - near-fault rupture directivity pulse

•Site - basin waves and basin edge waves



A Framing Equation

$$\lambda(\underline{DV}) = \int \int G(\underline{DV} \mid \underline{DM}) dG(\underline{DM} \mid \underline{IM}) d\lambda(\underline{IM})$$

- **DV** Decision Variable(s) (costs, lives lost, collapse limit states, ...)
- **DM** Damage Measure(s) (displacements, fractures, ...)
- **IM** Ground Motion Intensity Measure(s) (PGA, S_a , ...)





CHARACTERIZATION OF SEISMOGENIC SOURCES



SPECIFICATION OF RECURRENCE RELATIONSHIP





CALCULATION OF EXCEEDANCE USING ATTENUATION EQUATION



CALCULATION OF PROBABILISTIC SEISMIC HAZARD (HAZARD CURVE)





U.C. Berkeley Campus Combined Relationships for Rock Sites



U.C. Berkeley Campus 5%-damped Equal Hazard Response Spectra





Mw 7.0, 5km, Rock, AS



Ground Motion Prediction Models



Ground Motion Variability



GROUND MOTION VARIABILITY

For large earthquakes, site-to-site variability is much larger than event-to-event variability

So large earthquakes are basically similar, but a given earthquake looks different at different sites due to source, path and site effects that may in part be predictable

BROADBAND TIME HISTORY SIMULATION PROCEDURE

Elastodynamic Representation Theorem:

Ground motion U(t) can be calculated from the convolution of the slip time function D(t) on the fault with the Green's function G(t) for the appropriate distance and depth, integrated over the fault rupture surface

 $U(t) = \sum D(t) * G(t)$

Combine long period and short period simulations to generate broadband time history

EARTHQUAKE SOURCE REPRESENTATION

Kinematic model of shear dislocation spreading over the fault rupture surface, in which several processes are coherent at long periods (>~ 1 second) and less coherent at short periods:

- •Radiation Pattern
- •Rupture Velocity
- •Slip Velocity

Near fault pulse caused by rupture directivity and radiation pattern



SEISMIC WAVE PROPAGATION

Green's functions calculated for the required distance and depth ranges in a crustal structure model

Long periods - 1D: frequency wavenumber integration

- 3D: finite difference
- these methods are deterministic

Short periods - generalized rays

- empirically incorporate stochastic effects





Closest Distance (km)





Figure 1. Comparison of recorded (top row) and simulated (middle and bottom rows) displacement, velocity and acceleration time histories at Arleta from the 1994 Northridge earthquake, plotted on a common scale, with peak value given in the top left corner. Source: Somerville et al. (1995).



RUPTURE DIRECTIVITY EFFECT

- •Due to propagation of rupture toward a site
- •Large pulse of horizontal motion in the direction normal to fault strike
- •Large response spectral acceleration at periods longer than 0.5 second on fault normal cmpt.

1995 Kobe; 1994 Northridge





Figure 2. Relation between period of fault normal pulse and Mw for forward directivity



Magnitude Scaling Of Near Fault Ground Motions

- The forward directivity pulse is narrow band
- The period of the pulse increases with magnitude
- The pulse causes a peak in the acceleration response spectrum whose period increases with magnitude

Abrahamson and Silva, 1997, Strike-Slip, Soil, 5 km



Abrahamson and Silva, 1997, Strike-Slip, Soil, 5 km Somerville et al. (1997) with Directivity and Fault Normal





Abrahamson and Silva, 1997, Strike-Slip, Soil, 5 km Somerville (2001) with Directivity and Fault Normal

Mw 8.0	
Mw 7.5	
Mw 7.0	
Mw 6.5	



Rupture Directivity Models Mw 7.0, 5km, Soil









Buried and Surface Faulting Mw 7.0, 5km, Soil, AS



Turkey and Taiwan Spectra Mw 7.5, 5km, Soil, AS





Difference in Ground Motions: Buried vs. Surface Faulting

- Ground motions from buried faulting appear to be stronger than from surface faulting
- There may be differences in rupture dynamics between confined and runaway rupture









BASIN AND BASIN EDGE EFFECTS

•Due to trapping of body waves that enter a sedimentary basin through its thickening margins

•Cause increased amplitude and duration of ground motions with periods longer than 1 second

•Largest effects are located near the edges of basins due to the constructive interference of direct and diffracted arrivals

Kobe; Santa Monica (Northridge eq.)



Simulated Peak Velocity (0.1-0.8 Hz)

Log 10 Velocity (cm/sec)

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0.0	0.	4 0	6 0.	8 1.	0 1.:	2 1.	4 1.6	5 1.8	20	24



Vs (km/sec)

4.1

3.2

27 24 21

- 1.7

0.9



CONCLUSIONS

Variability in Ground Motions from Large Earthquakes:

•dominated by site-to-site variations that are partly predictable

Applications of Seismological Ground Motion Models:

- •provide more realistic representations of complex earthquake source, wave propagation, and site effects than simple empirical magnitude - distance - site category models
- •explain unusually large ground motions
- •reduce the uncertainty in ground motion prediction
- •provide time histories for use in performance based design

